

STARDUST IN THE CO3.0 CHONDRITE NWA 8631: LOW ABUNDANCE OF PRESOLAR SILICATES.

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Introduction: Only a handful of primitive meteorites are characterized by high presolar silicate abundances [1], and CO3.0 chondrites (e.g., ALHA 77307, LAP 031117 and DOM 08006), in particular, are highly unequilibrated and minimally altered carbonaceous chondrites that contain among the highest presolar silicate abundances (150–200 ppm) found in meteorites [2–4]. Like type 3 ordinary chondrites, these meteorites are classified into a metamorphic sequence from type 3.0 (e.g. ALHA77307) to 3.8 (e.g., Isna) based on their mean Cr contents in ferroan olivine from type II chondrules [5, 6]. Recently, a new possible CO3.0 chondrite (a 37-gram single stone), Northwest Africa 8631 meteorite (NWA 8631), was found in the Sahara Desert in Northern Africa. Initial characterization showed that ferroan chondrule olivines in NWA 8631 have a mean Cr₂O₃ content (0.35 ± 0.06 wt.% [7]) that is similar to that in the highly primitive CO3.0 chondrite, DOM 08006, suggesting that it is also a type 3.0 CO chondrite. However, it was also noted that both the matrix abundance (~53%) and average chondrule size (240 ± 130 μm) are larger than usual for CO chondrites [7]. Here, we report preliminary data on the presolar grain abundances in this meteorite suggesting that it may not be a type 3.0 CO chondrite.

Sample and Experimental Methods: We obtained a polished section of NWA 8631 from the meteorite collection at the European centre research and teaching in environmental geosciences (CEREGE; Aix-en-Provence, France). We then carried out raster ion imaging of carbon and oxygen isotopes in fine-grained matrix areas in NWA 8631 using the NanoSIMS 50 at Washington University. A focused Cs⁺ primary beam of ~1 pA (~100 nm in diameter) was rastered over areas of 10×10 μm². Secondary ions of ^{12,13}C⁻ and ^{16,17,18}O⁻, as well as secondary electrons, were simultaneously acquired in multicollection mode. We measured a total area of 8,400 μm².

Results and Discussion: To date, we have identified a total of nine presolar grains in NWA 8631, including six O-rich presolar grains (five ¹⁷O-rich grains and one ¹⁸O-rich grain) and three C-anomalous grains. Based on the classification proposed by [8], four grains belong to Group 1, with excesses in ¹⁷O and sub-solar to solar ¹⁸O/¹⁶O ratios, and are believed to have originated from the envelopes of low- to intermediate-mass AGB stars of close-to-solar metallicity. Another grain is characterized by an extremely high ¹⁷O/¹⁶O ratio (¹⁷O/¹⁶O = 7.4 ± 0.1 × 10⁻⁴), higher than the upper limits (¹⁷O/¹⁶O = 6 × 10⁻³) predicted by model calculations of the first and second dredge-up in low- and intermediate-mass AGB stars [9]. These Extreme Group 1 grains may originate from ONe nova [10]. However, all nova models also predict very large depletions in ¹⁸O that are not observed in most nova oxide-silicate grains, suggesting mixing of the nova ejecta with solar composition material. Finally, the last grain is characterized by an excess in ¹⁸O relative to solar (Group 4) and may have formed in the ejecta of a Type II supernova [11].

The three C-anomalous grains are all enriched in ¹³C relative to solar and are likely SiC grains. One grain is likely a mainstream SiC (¹²C/¹³C = 53.8 ± 0.4), with an origin in a low-mass AGB stars of close-to-solar metallicity. The two other grains are characterized by ¹²C/¹³C ratios between 12–16 and are possible AB SiC grains, whose origin(s) are still unclear, with born-again AGB stars, J-type stars, Type II supernovae or novae as possible stellar sources [12].

Based on the nine presolar grains identified, we calculated O- and C-anomalous grain abundances of 99⁺³¹⁵₋₅₇ ppm and 83⁺²⁵⁰₋₆ ppm, respectively. While our data are still limited, the abundance of O-rich grains is significantly lower than what has been observed in other CO3.0 chondrites [2–4], indicating that this meteorite might have experienced more secondary processing (aqueous alteration or thermal metamorphism). Our preliminary observation thus suggests that, despite its initial classification, NWA 8631 is likely not a type 3.00 CO chondrite but is of higher petrologic type (possibly 3.1 or 3.2). We will continue our NanoSIMS search for presolar grains in this meteorite to better characterize its presolar grain abundances. Using Auger spectroscopy, we will also determine the elemental compositions of the presolar grains that we identified to investigate any further evidence of secondary processing.

References: [1] Floss C. and Haenecour P. 2016. *Geochemical Journal* 50:3–25. [2] Bose et al., 2012. *Geochimica et Cosmochimica Acta* 93, 77–101. [3] Haenecour et al., 2015. *Lunar Planetary Science Conference XLVI*, #1160. [4] Nittler et al., 2013. *Lunar Planetary Science Conference*, #2367. [5] Grossman and Brearley, 2005. *Meteoritics & Planetary Science* 40, 87–122. [6] Davidson et al., 2014. *Geochimica et Cosmochimica Acta* 139, 248–266. [7] Meteoritical Bulletin Database (<http://www.lpi.usra.edu/meteor/>). [8] Nittler et al., 1997. *The Astrophysical Journal* 483, 475–495. [9] Gyngard et al., 2011. *Lunar and Planetary Science Conference XLII*, #2675. [10] Leitner et al., 2012. *The Astrophysical Journal Letters* 754, L41 (46pp). [11] Nittler et al., 2008. *The Astrophysical Journal* 682, 1450–1478. [12] Amari et al., 2001. *The Astrophysical Journal*, 559: 463–483.

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