

HAMMADAH AL HAMRA 193: AN AMPHIBOLE-BEARING WINONAITE. C. Floss, B. Jolliff, J. Reid, G. Benedix; Dept. of Earth and Planetary Sciences, Washington University, St. Louis, MO 63130, USA; (email: floss@wuphys.wustl.edu).

Introduction: Winonaites are reduced meteorites with oxygen isotopic compositions distinct from those of other primitive achondrites, except for silicate inclusions from IAB irons [1,2]. They have equigranular textures and chondritic mineralogies, but can contain mm- to cm-sized regions with different textures and/or mineralogies. These include plagioclase-diopside-rich regions, coarse-grained olivine-rich areas and/or large poikilitic calcic pyroxenes [2]. Here we report on a texturally and mineralogically unusual winonaite, Hammadah al Hamra (HaH) 193.

Results: HaH 193 was found in 1996 in the Libyan desert as a single stone of 259 g. Its identification as a winonaite [3] was confirmed by oxygen isotopic analysis ($\delta^{18}\text{O} = +4.79$, $\delta^{17}\text{O} = +1.90$; R. Clayton, pers. commun.). We obtained a thin section of HaH 193 for a trace element study of winonaites [4] and discuss here its mineralogy and petrology.

Our section of HaH 193 is dominated by very large (up to 5 mm) optically continuous orthopyroxenes that poikilitically enclose smaller grains of olivine, plagioclase and occasionally clinopyroxene. Two distinct lithologies occur interstitial to the poikilitic orthopyroxenes. One consists of clumps of equigranular 200-300 μm olivine grains, similar to areas described in Winona and other winonaites [2]. The other lithology contains abundant amphibole and clinopyroxene, as well as olivine and lesser amounts of plagioclase and orthopyroxene. The amphiboles are large (up to 2 mm) and poikilitically enclose smaller grains of primarily olivine, but also plagioclase, clinopyroxene and orthopyroxene.

Major element mineral compositions are similar to those seen in other winonaites (orthopyroxene: $\text{Fs}_{4.9-5.1}\text{Wo}_{1.7-2.0}$; clinopyroxene: $\text{Fs}_{1.3-2.2}\text{Wo}_{43.2-44.2}$; olivine: $\text{Fa}_{4.3-4.6}$; plagioclase: $\text{An}_{20.9-21.4}$). The amphibole in HaH 193 is nearly stoichiometric fluoro-edenite, $\text{K}_{0.1}\text{Na}_{1.1}\text{Ca}_{1.7}\text{Mg}_{4.8}\text{Fe}_{0.1}\text{Ti}_{0.1}\text{Si}_{7.0}\text{Al}_{1.0}\text{O}_{22}(\text{F})_{1.8}$. Raman spectroscopy confirms the amphibole structure and shows no OH peak. Edenite is rare and endmember fluoro-edenite has only recently been identified terrestrially [5]. This is the first reported meteoritic occurrence.

Discussion: Textures suggest that fluoro-edenite is replacing clinopyroxene in HaH 193, probably in a reaction between clinopyroxene and plagioclase. This is supported by trace element evidence. Clinopyroxene and edenite have similar bow-shaped REE patterns with negative Eu anomalies. However, relative to clinopyroxene, the edenite has lower trivalent REE abundances ($\sim 2x$), and higher abundances of Eu ($\sim 3x$), Ba ($\sim 15x$), Na ($\sim 10x$) and K ($\sim 100x$), all elements that are typically enriched in plagioclase. The source of the F is probably apatite. Apatite is rare in HaH 193, but several grains associated with edenite have F contents ranging from 3.9 to 4.2 wt.%, in contrast to the chlorapatite commonly found in other winonaites.

Although HaH 193 is fairly weathered and our section contains several alteration-filled veins, the edenite appears to be of pre-terrestrial origin. Evidence for this includes textural relationships and its OH-free composition, as well as the fact that HaH 193 minerals show little evidence of the terrestrial alteration commonly observed in other hot desert meteorites [*e.g.*, 6].

References: [1] Clayton and Mayeda (1996) GCA 60, 1999. [2] Benedix et al. (1998) GCA 62, 2535. [3] Grossman (1998) MAPS 33, A221. [4] Floss et al. (2003) this vol. [5] Gianfagna and Oberti (2001) Am. Mineral. 86, 1489. [6] Crozaz et al. (2002) GCA 66, A158.